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RESEARCH ARTICLE



ASSESSING THE ACID MINE DRAINAGE POTENTIAL OF A STOCKPILED GOLD BEARING SULPHIDE ORE

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ARTICLE DETAILS	ABSTRACT
<i>Article History:</i> Received 01 June 2021 Accepted 03 July 2021 Available online 29 July 2021	This work seeks to assess the acid mine drainage (AMD) potential of stockpiled sulphidic gold bearing ore that was left untreated for a period of four (4) years. The mineralogical composition of the stockpiled sulphidic
	gold ore was determined using X-Ray Diffractometry (XRD) and Reflected Light Microscope. The AMD potential and kinetic studies on the ore sample were determined using Acid Base Accounting (ABA) and Humidity Cell (ASTMD 5744). The XRD and microscope analyses showed that the gold ore contain arsenopyrite and pyrite as the major sulphide minerals. The ABA confirmed that the gold sulphide ore could be a potential source of AMD with respect to AMD marginal of safety if the ore is left untreated for a very long period. Humidity Cell kinetic analysis showed a possibility of mobilizing Arsenic (As) into the geo-environment when there is a long period of oxidation of abandoned sulphide run-of-mine (ROM) ore.
	KEYWORDS
	Acid Pass Accounting Argonic Sulphate Concentration Heavy Metals Concentration Conductivity

Acid Base Accounting, Arsenic, Sulphate Concentration, Heavy Metals Concentration, Conductivity

1. INTRODUCTION

Natural geochemical processes that oxidize metal sulphides exposed at the earth's surface by anthropogenic activities such as mining, results in acid mine drainage (AMD). AMD occurs when mine rocks or mine waste rocks and tailings containing sulphide-bearing minerals such as pyrite, arsenopyrite, chalcopyrite and pyrrhotite are exposed to molecular oxygen-rich environment and water (Warhurst and Noronha, 2000; Simate and Ndlovu, 2014). Iron and sulphur oxidizing bacteria are also known to catalyze these reactions at low pH thereby increasing the rate of reaction by several orders of magnitude (Nordstrom and Southam, 1997). These reactions however, occur naturally but at a rather slow rate. The reaction of pyrite with oxygen and water produces a solution of ferrous ion (Fe2+), sulphate ion (SO42-) and protons (H+ acid) (Jennings et al, 2008; Simate and Ndlovu, 2014). The sulphur oxidation produces the sulphuric acid and the ferrous ion is further oxidized to ferric ion, which forms the orange ferric hydroxide precipitate (Fe (OH)₃) as summarized in Equation.1. As at today, AMD has been recognized as one of the main environmental challenges derived from mining activities (Dold and Fontbote, 2001; Holmstrom et al., 2001; Romero et al., 2010; Naidu et al., 2019). The resultant acidic leachate may be transported to the surrounding environment and contaminate soils, sediments, ground and surface waters (Bain et al., 2000; Armienta et al., 2001; Jung, 2001; Romero et al., 2010; Dong et al., 2018; Skousen et al., 2019). Often, the chemical characteristics associated with AMD are low pH values (high acidic water), high concentration of metals (iron, aluminium and manganese and other heavy metals) and metalloids (of which arsenic is generally of great concern) and sulphate ions in solution (Johnson and Hallberg, 2005; Mendez-Ortiz et al., 2007).

 $4\text{FeS}_2 + 150_2 + 14\text{H}_20 \rightarrow 4\text{Fe}(0\text{H})_3 + 8\text{S}0_4^{2-} + 16\text{H}^+$ (1)

Due to the above stated detrimental effects, AMD generation associated with sulphide bearing waste rocks and tailing material have been studied in detail over the past years (Dold, 2010; 2014; Banerjee, 2014; Simate and Ndlovu, 2014; Celebi and Öncel, 2016; Zhou et al., 2017; Fosu et al., 2020). Kuma (2003) and Asamoah et al. (2007) reported the occurence of AMD in the Tarkwa mining area in Ghana. Ndur et al. (2010) assessed AMD potential of the Coral Snake Waste Rock Dump located close to the Enkansu and Kaw streams in Obuasi. Ghana. With the above notwithstanding, little has been done on the AMD potential of sulphide bearing Run of Mine (ROM) ore which stays on the ROM pad/ Stockpile for several years and interact chemically with oxygen and rain water. For instance, Brierley and Brierley (2001) reported that 1.1 million dry-tons of about 80% pyrite of cobaltiferrous pyrite concentrate (grading 1.38 % cobalt) produced from the Kilembe mine had been stockpiled for over 30 years before the cobalt recovery. Also, the Escondida Low-Grade Sulphide bioleach project began after the ore had been stockpiled for 2 years (Clark et al, 2006).

Storing these sulphide bearing ores for such long periods before the beneficiation process causes the oxidation of unstable sulphides (Wills, 1997) resulting in AMD. Impoundments have been categorized as main sources of AMD although open pit high walls, underground workings and ore stockpiles contribute significant volumes of AMD (US-EPA, 2005; Romero *et al*, 2010). As stated by Asare-Donkor *et al* (2015), the Anglogold Ashanti Mine, Obuasi which was shut down from 2014 to 2019 (Ghanaian times, 2020) had large deposits of mine wastes, ore stockpiles and waste rocks that have become a heap around the extraction plants. Also, illegal miners in Ghana strip about 28 million tons

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of ore annually (Agbesi, 2017). In 2017, activities of these miners were brought to a halt by a directive from the government due to environmental issues. Hence, resulting in abandoned stockpiles of ROM ore in most illegal mining areas in Ghana for over two years. It is therefore important to ascertain the mineralogy of ROM ore and assess its AMD potential. This work attempts to investigate the AMD potential of sulphide rich ROM ore at Obuasi Ghana. Findings from the investigations will assist in managing seepage from the ROM/stockpile area.

1.1 Geology of the Study Area

The geology of Ghana can be divided into four distinct geotectonic units namely: the West African Craton, Platform, Pan-African mobile belt and coastal basins. The West African Craton was stabilised towards the end of the Eburnean Orogeny at approximately 2 Ga (Leube et al., 1990) and comprises the Paleoproterozoic Birimain Metasedimentary and Metavolcanic rocks which have been intruded by basin and belt type granitoids respectively. Overlying the Birimain Metavolcanic belt is the Tarkwaian Group which is believed to be of shallow water continental origin derived from the Birimain and its associated granitoids. The Platform area is dominated by the Neoproterozoic to Paleozoic Voltaian basement complex which is from the base made up of Kwahu (Morago) group, Oti-Pandjari group and Obosum group. The Pan-African Mobile belt is also made up of Buem, Togo and Dahomeyan structural units. The Buem structural unit consists of a thick, lower sequence of clastic sediments with some carbonate and tillite units succeeded by clastics and volcanics that include mafic flow units and pyroclastics (Kesse, 1985). Also, the highly metamorphosed Togo structural unit is predominantly composed of quartzite, chert, mica schist and phyllite while the Dahomeyan is predominantly gneissic in composition (Duodu et al., 2009). The coastal basin comprises the Phanerozoic units are also made Paleozoic coastal sediments, Mesozoic -Tertiary sediments and basins and Mesozoic intrusives.



Figure 1: Geological map of Ghana showing the study area (modified after Davis et al. 1994)

The study area (Obuasi) falls within the Birimian of Ghana, precisely at the contact between the Ashanti belt and Kumasi basin (Fig.1) where gold-bearing quartz veins and sulphides are the two dominant ore types (Fougerouse et al., 2013). The area is underlain by Paleoproterozoic Birimian Metasedimentary (argillites, carbonaceous schist, wackes and chemical sediments), Metavolcanic rocks (basic to intermediate volcanic rocks) and the Tarkwaian sediments to the SE (Osae et al. 1999). Deformation within the area is marked by the carbonaceous shear zones that follow the general strike of the Birimian rocks (ie. NE). The shear zone is characterized by different generations of quartz veins (smoky blue-gray quartz veins, ribboned milky white quartz veins and quartzcalcite veins) which are usually mined as stockwork. Sericitization, carbonitization and sulfidation are predominant alterations in the area (Osae et al. 1999).

2. MATERIALS AND METHODS

2.1 Sampling

Run of Mine (ROM) ore samples (predominantly carbonaceous schist and quartz veins) were collected from six different locations of a stockpile. The samples were homogenized with a shovel and then further sampled using coning and quartering method to obtain 5 kg of the ore material. The prepared sample was then preserved at a temperature of 4 °C for further analysis.

2.2 Mineralogical Investigations

The mineralogy of the ROM ore was determined by using physical observation, reflected-light microscopy (RLM) and X-ray diffractometry (XRD). Before bulk characterization, the sample was washed with distilled water to get rid of fine particles sticking on the rock surfaces. Polished mounts (< 45 μ m) were prepared for RLM analysis using conducting graphite doped epoxy. The samples for the XRD analysis were pulverized and processed by placing on a powder plate and compacting with a glass slide. The XRD powder patterns (scans) were obtained using a Rigaku D/MAX 2500 rotating anode powder diffractometer using the JADE version 9.0 coupled with the ICSD and ICDD diffraction databases. The measurement was done with two theta angular range of 5 – 70°, step size of 0.020°, scan speed of 1°/min, voltage of 50 kV and current of 260 mA.

2.3 Acid Base Accounting (ABA)

The most commonly and dependable static test used to predict AMD from most sulphide rock-bearing mineral is the standard acid-base accounting (ABA) (Yeheyis et al., 2009; Skousen, 2017). The ABA test determines the balance between maximum acid-generating processes (oxidation of sulphide minerals) and acid-neutralising processes (dissolution of alkaline carbonates, displacement of exchangeable bases and weathering of silicates) (Sobek et al., 1978; Ritcey, 1989, Skounsen et al., 2002; Yeheyis et al., 2009; Banerjee, 2014; Yucel and Baba, 2016). Acid Base Accounting (ABA) of the ROM ore samples was carried out using the MEND method (1991). In this experiment, modified Sobek method was employed in the ABA analysis. Results of the experiment were computed using Equations (2), (3) and (4). Finely ground ROM ore sample (2 g ;< 45 μ m) was soaked in 40 ml 0.1 N HCl for 24 hrs and then the resulting solution was titrated with 0.1 N NaOH until a pH of 7 was obtained. The sample was also analyzed for total sulphur content using the LECO sulphide analyzer.

$$NP = \frac{50a \left[x - {\binom{b}{a}y}\right]}{c}$$
(2)

 $MPA = 31.25 \times \% S_{sulpide}$ (3)

$$NNP = NP - MPA$$
(4)

where: *NP* is the Neutralization Potential (kg CaCO₃/t), *MPA* is the maximum potential acidity (kg CaCO₃/t), NNP is the Neutralization Potential (kg CaCO₃/t), 'a' is the normality of HCl, 'b' is the normality of NaOH, 'x' is the volume of HCl (mL), 'y' is the volume NaOH (mL) and *S* is the sulphur content of the sulphide.

2.4 Laboratory Weathering of ROM Ore

The ROM ore was subjected to standardized laboratory weathering analysis using the humidity cell test procedure (ASTM D 5744 – 07; Anon, 2007). This was to enhance the study of the kinetics involved in the rate at which weathering product mass (SO_4^{2-}) is released from a known mass (250 g) of the ROM ore when it is leached every week using a fixed volume (500 mL) of de-ionised water. About 2 kg of the material was homogenized; hand crushed using a laboratory mortar and pestle and sieved to obtain a particle size range of 80 % passing a 4.76 mm aperture size sieve. The sample was then thoroughly washed several times with distilled water until the conductivity of the spent rinsing water was almost the same as that of the fresh distilled water. The residue was then dried with paper towels and further sampled using a Jones riffle to obtain

two sets of 250 g of ROM ore. The material (250 g) was transferred into a humidity cell column of height, internal diameter and wall thickness of 20.3 cm, 10.2 cm and 0.635 cm respectively lined with a geotextile (Fig. 2). The experiment was conducted for ten weeks at an average temperature and humidity of 22 ± 1 °C and 21 % RH respectively. An EXTECH 45160 was used for the temperature and humidity measurements. The setup was supplied with air using a standing table fan. Sulphate analysis on resulting weekly leachate solutions was obtained using the Waters High Pressure Liquid Chromatogram (HPLC) with a flow rate of 1.2 ml/min. The column and detector types were IC Pack anion 4.6 x 50 mm and Waters 432 Conductivity Detector respectively. The resulting weekly leachates were also analyzed for selected heavy metals using inductively coupled plasma-optical emission spectroscopy (ICP-OES).



Figure 2: Schematic set-up for weathering of ROM ore in Humidity Cell

3. RESULTS AND DISCUSSION

3.1 Mineralogical Investigations

Macro and microscopic examinations of hand specimen and polished sections as shown in Fig. 3 and Fig. 4 respectively revealed that the ROM ore contained arsenopyrite and pyrite as major sulphides minerals. Fig. 3a is a highly sheared and foliated mineralized carbonaceous schist with deformed pyrite. Fig. 3b is a typical texture of "classic" crush zone of finely disseminated grains of deformed pyrite and needle-like arsenopyrite. Figs. 3c1 and 2 are grains of arsenopyrite and pyrite with quartz vein and microfracture respectively. The size of the sulphides in the ROM ore was between 100 and 3000 µm. The sulphide grains were observed to be sub perpendicular to foliation in some sections of the ore. In some instances, finely disseminated grains were observed to be occurring along microfractures and shear planes. Arsenopyrite grains in the ROM ore were idiomorphic crystals with a rhomboidal shape as shown in Fig. 4a. They had well-defined straight edges and appeared whitish. Also, some large crystals of arsenopyrite in the ROM ore were observed to be massively fractured. Other arsenopyrite grains, however, appeared elongated with both rough and smooth edges (Fig. 4b). Their sizes ranged from 200 to 1200 µm with few fractures.

Pyrite in the ROM ore appeared as brassy yellow to yellowish-white in air when viewed using a microscope. The sizes range from 40 to 2000 μ m in a range of textures and crystal forms. The Pyrite grains appeared predominantly as spongy or porous masses with intergrown material as well as framboids at the rough ends of some of the crystals (Fig. 4a). Qualitatively, few framboids occurred as discrete particles. Most of the framboids majorly occurred as aggregates but are found at the edges of the pyrite crystals. The pyrites are more deformed compared to arsenopyrite due their spongy nature.

The X-ray diffractograms (XRD, Fig. 5) showed that the ROM ore was made up of quartz, illite, dolomite, chlorite-serpentine, albite, muscovite, paragonite, rutile, pyrite, arsenopyrite and pyrrhotite in order of decreasing abundance. The average composition of the main ore types at the mines are made up of about 35% quartz and 60 – 70% sulphide material. The run of mine taken for this analysis has key geochemical compositions in relative proportions such as: Fe (45.3%); As (32.6%); S (20.0%) and other base metals constitute about 2.1% dry weight (Oberthur *et al.*, 1994; Osae *et al.*, 1995; Foli *et al.*, 2015).



Figure 3: Hand specimen of ROM ore showing (a) a Highly sheared and foliated Mineralised carbonaceous schist (b) a Typical Texture of "Classic" Crush Zone of Finely Disseminated grains of Pyrite (Py) and Arsenopyrite (Aspy) (c1) Quartz vein within the Mineralised carbonaceous schist (c2) Pyrite (Py) along Microfractures and Shear Planes



Figure 4: Photomicrographs of polished sections of rhomboidal arsenopyrite (Aspy) and deformed pyrite (py) crystals b) elongated grain of arsenopyrite



Figure 5: X-ray Diffractogram of Bulk Material Mineralogy (ROM Ore) showing the presence of quartz, illite, dolomite, chlorite-serpentine, albite, muscovite, paragonite, rutile, pyrite, arsenopyrite and pyrrhotite in order of decreasing abundance.

3.2 Acid Base Accounting

The result of acid base accounting (ABA) on the ROM computed using Eqns (2), (3) and (4) is presented in Table. 1. ABA was conducted to help assess the acid producing and acid neutralising potential of the ROM ore prior to mineral processing. The calculated maximum potential acidity (MPA) of the ROM ore was 30.53 kg CaCO₃t⁻¹ while the neutralisation potential (NP) is 39.52 kg CaCO₃t⁻¹. The MPA value is attributed to the presence of few reactive sulphide minerals (pyrite, arsenopyrite and pyrrhotite). Again, the maximum NP value could be attributed to the following acid neutralising minerals such as slightly soluble dolomite, intermediate weathering chlorite-serpentine, very slow weathering muscovite (Sverdrup, 1990). The net neutralization potential (NNP) was found to be 8.99 kg CaCO₃t⁻¹. This means that there is enough NP (carbonate = calcite) to maintain the ROM ore neutral should it become chemically reactive leading to AMD. The neutralisation potential ratio (NPR) defined as (NP/MPA) was also found to be 1.2943. One of the criteria used to determine the acid producing potential of a rock-bearing mineral is the NNP value (Lottermoser, 2010). Conceptually, rocks with positive NNP values have no potential to generate acid, whereas rocks with negative NNP values do. A factor of safety is used in practice and rocks with significantly positive NNP values are generally considered as having no potential to generate acid (> $+20 \text{ kg CaCO}_3t^{-1}$); rocks with a significant negative NNP value (< -20 kg CaCO₃t⁻¹) are potentially acid generating (Yucel and Baba, 2016). Rock-bearing sulphide minerals with intermediate NNP values have uncertain acid generation potential (-20 kg CaCO₃t¹ < NNP < +20 kg CaCO₃t¹) (Ferguson and Morin, 1991; Miller et al., 1991; Lottermoser, 2010, Yucel and Baba, 2016).

The ROM ore as examined for AMD potential in this study can hence be said to be uncertain since the value of $8.99 \text{ kg } \text{CaCO}_3\text{t}^1$ falls in the uncertain range and this can be investigated using kinetic leaching studies to establish the domains. The NPR also indicates the relative margin of safety within a rock-bearing mineral for AMD. If NPR value is less than 1, the sample is said to be acid generating and if the NPR is between 1 and 2 (1<NPR<2), the sample is considered to be potentially acid generating, if the NPR is 2 or more, the indication is that the rock-

bearing mineral would have a higher probability to remain near-neutral and should not be problematic with respect to AMD (Price *et al.*, 1997; Price and Errington, 1998; Skousen et al., 2002; Yucel and Baba, 2016). Again, according to Brodie et al. (1991), NPR less than 1 of a sample is said to be acid generating, NPR value of sample between 1 and 3 is said to be uncertain and NPR greater than 3 is said to be non-acid generating. Using the NPR marginal safety condition with respect to AMD, then ROM ore is said to have a possible AMD potential since a value of 1.295 falls between NPR values 1 and 2 according to Price *et al.* (1997), and uncertain since a value of 1.295 falls between NPR values 1 and 3 according to Brodie *et al.* (1991).

Table 1: Summary of Acid Base Accounting (ABA) Results						
ABA Parameter	ROM					
Paste pH	7.58					
Total Sulphur (%)	0.977					
Neutralisation Potential (NP) kg CaCO $_3$ t 1	39.52					
Maximum Potential Acidity (MPA) kg CaCO3t ⁻¹	30.53					
Net Neutralisation Potential (NNP) in kg CaCO3t-1	8.99					
Neutralisation Potential Ratio (NPR)	1.295					

3.3 Weathering of ROM ore Results

The results of the leaching (weathering) test conducted on the ROM indicate that there is averagely low heavy metal concentration release from the ROM ore with slightly acidic to neutral pH (6.64 - 7.98) solution as shown in Table.2. An average oxidation-reduction potential (Eh) value of about 600 ± 63 mV indicates moderate oxidation of the ROM ore sample during the leaching process which allowed moderate reactivity and smaller electrical conductivity in the leachate. Ferguson and Morin

(1991); Patterson and Ferguson (1994) however established that rocks with NNP greater than 0 kg CaCO₃t¹ do not produce acid. This confirms the leaching test since the NNP value of the ROM is 8.99 kg CaCO₃t⁻¹. If the sample is strongly generating acid then the pH of the leachate solution is expected to be < 6 with high or significant concentrations of metal contaminants. The low metal concentration is also confirmed by the low average electrical conductivity (21 ± 9 µS/cm) of the leachate (Table. 2). If ROM ore sample is generating acid strongly, electrical conductivity levels between 1000 to 20000 µS/cm are expected (Sephton and Webb, 2017). Notwithstanding that, the ROM ore could have a low AMD potential but AMD can occur over long exposure periods of interaction between the ore and water in an atmosphere if the ore is left untreated for a very long time since there is a slight potential for AMD. The AMD produced will depend on the size of the exposed surface area of the type and quantity of sulphide minerals present as well as the host rock composition (Ziemkiewcz et al., 2003).

From Fig. 6, it can be seen that the release of sulphate from the leaching process increased with increasing time with a slow-release rate of approximately 0.423 mgL⁻¹day⁻¹ for every 250 g of the ore under the test conditions. AMD is characterized by a high sulphate concentration in the range of 500 - 10, 000 mg/L (Hill, 1968). In extrapolating for the limit of the sulphate ion concentration, it will take approximately 3 years for a sulphate ion concentration of 500 mg/L to be produced and 65 years for sulphate ion concentration of 10, 000 mg/L to be produced in the leachate. From Table. 2, concentrations of Fe, Pb, Zn and Cu were within the effluent permissible limits set by the Ghana EPA (2007) and (2010). However, concentration of Arsenic (As) was slightly above Ghana EPA (2007) and (2010) permissible limit for the 56^{th} day. This indicates that sulphide contained in stockpiled ores must be properly managed to prevent contamination of the geo-environment. The low concentration of Fe and As release may be due to the fact that the leachate samples had pH higher than 6. This is because low pH (< 4) accelerates the acid generation, dissolving Fe oxyhydroxides and releasing As (Andrade et al., 2008). Again, the presence of carbonate in the form of dolomite as shown in Fig. 5 could reduce the oxidation rates and Arsenic (As) mobilisation as seen in the work of Andrade et al. (2008). All these can be attributed to low oxidation rate of the sulphide, since the amount of leached sulphate was also low in the leachate samples (Fig. 6). This is in agreement with the authors (Nicholson et al., 1990; Jones et al., 2003) who reported that the rate of sulphide oxidation is reduced by increasing pH of the medium. Also, there could be retention of As by iron (hydroxides) which precipitate just after sulphide oxidation at high pH (> 4) (Pierce and Moore, 1982; Sun and Doner, 1998; Ladeira and Ciminelli, 2004; Gupta et al., 2005).

Figure 6: Plot of Cummulative sulphate concentration release from ROM Ore in days during a laboratory weathering analysis.

Table 2: pH, mean concentrations, conductivity and redox potential of metal species in leachate from laboratory weathering analysis									
Time		Heavy Metal Concentration (mg/L)						Eh vs. SHE	
Day	рН	Fe	As	Cu	Pb	Zn	(µS/cm)	mV	
0	6.64	0.07	0.02	0.01	0.01	0.03	38	622	
7	6.94	0.09	0.02	0.01	0.01	0.02	39	620	
14	7.10	0.15	0.02	0.00	0.01	0.01	25	648	
21	6.94	0.07	0.02	0.00	0.01	0.02	18	649	
28	7.63	0.04	0.01	0.00	0.01	0.02	20	722	
35	7.05	0.04	0.01	0.00	0.01	0.02	15	649	
42	7.55	0.05	0.02	0.00	0.01	0.02	17	557	
49	6.89	0.08	0.01	0.00	0.01	0.01	13	548	
56	7.98	0.10	0.22	0.00	0.01	0.02	15	519	
63	7.55	0.04	0.01	0.00	0.00	0.02	16	542	
70	7.39	0.07	0.01	0.00	0.00	0.01	15	522	
Ghana EPA Effluent Permissible Limit	6.0 - 9.0	2.0	0.2	1.0	0.1	2.0	750		

4. CONCLUSIONS

The net neutralisation potential (NNP) and neutralisation potential ratio (NPR) of the acid base accounting (ABA) were used to assess the possible AMD potential of stockpiled gold-bearing sulphide ore. The mineralogical analysis carried out on the gold ore showed that arsenopyrite and pyrite are the most common sulphide minerals present in the stockpiled material (predominantly composed of carbonaceous schist and quartz veins). Quartz, illite, dolomite, chlorite-serpentine, albite, muscovite, paragonite and rutile were the non-sulphide gangue minerals identified in the ROM ore. Some of these non-sulphide minerals could play a major or a minor role in the acid neutralisation due to their differences in relative reactivity.

The maximum potential acidity (MPA), neutralisation potential (NP), NNP and NPR associated with the ROM ore were $30.53 \text{ kg} \text{ CaCO}_3 t^{-1}$, $39.52 \text{ kg} \text{ CaCO}_3 t^{-1}$, $8.99 \text{ kg} \text{ CaCO}_3 t^{-1}$ and 1.295 respectively. Considering the ROM ore sample, the NP exceeded the MPA, indicating that the ore sample generally has a low AMD potential or uncertain. Decisively, the pH from the weathering test were slightly acidic to neutral pH values. Very low metal concentrations which were uncharacteristic of a strong acid generating material were obtained. The pH condition, slow release and low concentration of sulphate suggest low oxidation of sulphide mineral in the ROM ore per the conditions provided in the leaching test work.

There is a possibility of mobilizing Arsenic (As) into the geo-environment by ROM ore oxidation since the concentrations slightly increased above the Ghana EPA permissible limit on particular day. Subsequently, AMD could occur over a long period of exposure of interaction between the ROM ore and water under right atmospheric conditions if the ore is left untreated or abandoned for a very long period.

5. Recommendations

In order to expand the frontiiers of knowledge in potential AMD pollution of sulphidic run of mine ore, it is strongly recommended that all other AMD indicators in the ore be ascertained by prolonging the study time for six plus years.

Also kinetic net acid generation tests must be developed to provide a qualitative indication of the lag to onset of AMD from the sample.

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